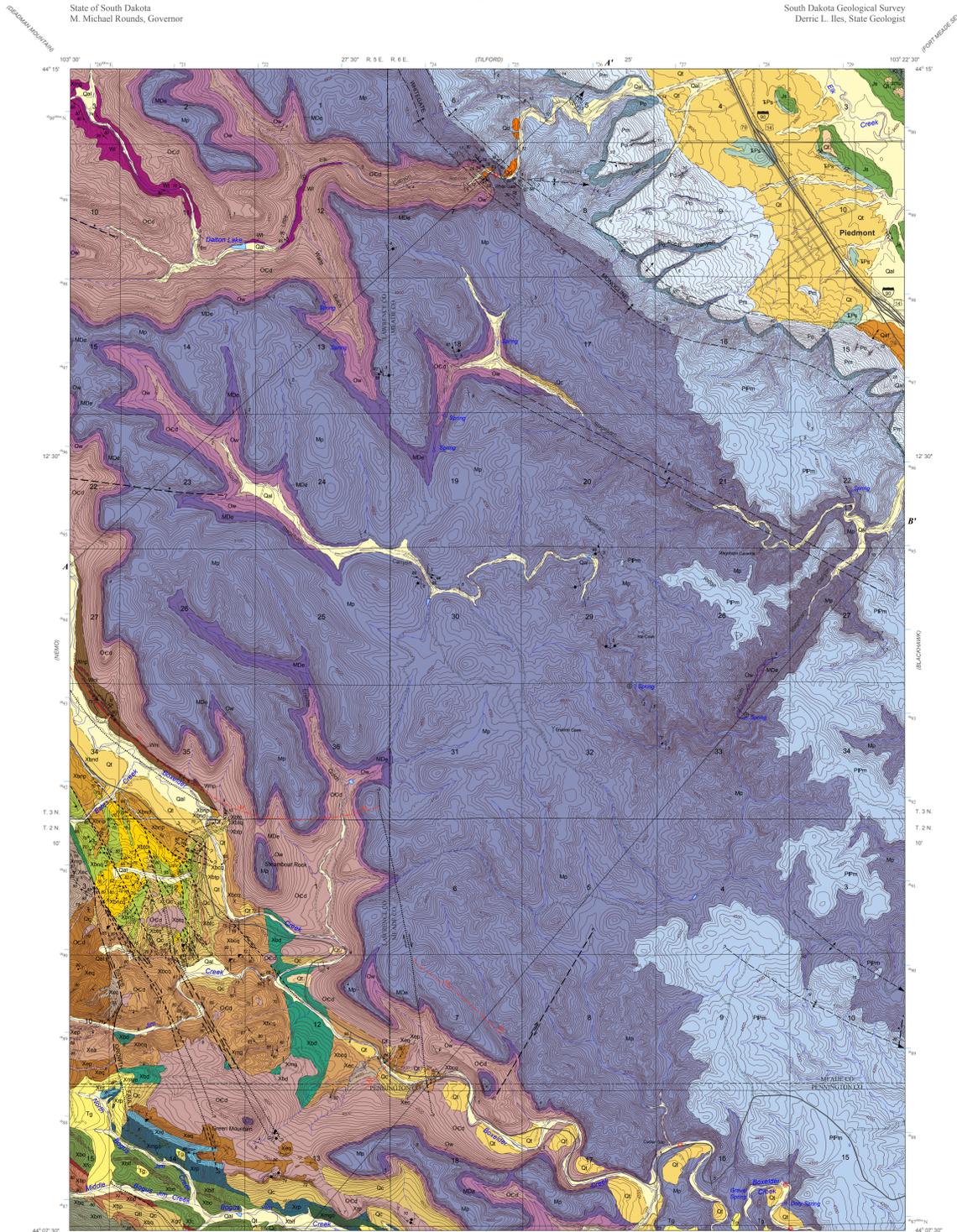


# Geologic Map of the Piedmont Quadrangle, South Dakota

J.A. Redden  
2009

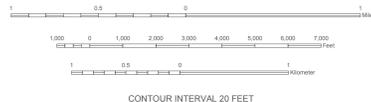
Prepared in cooperation with the South Dakota School of Mines and Technology



Map base modified from U.S. Geological Survey 1:24,000-scale Piedmont digital line graph.  
Projection is Universal Transverse Mercator, Zone 13.  
Datum is 1983 North American.

The Geological Survey, Department of Environment and Natural Resources, engages in an ongoing data collection and interpretation process. An outcome of that process is to reflect those interpretations on maps such as this one. Reasonable efforts have been made to ensure that this map accurately reflects the source data used in its preparation. This map is date specific. As additional data become available, geologic interpretations may be revised and the map may be updated by the Geological Survey. This map should not be enlarged or otherwise used in an attempt to interpret more detail than can be seen at the 1:24,000 scale.

SCALE 1:24,000



NORTH



Edited by M.D. Fahrenbach  
Digital Cartography by L.L. Roenstad, E.S. Beck,  
M.D. Fahrenbach, B.A. Fagnan, and L.D. Schulz

South Dakota  
Quadrangle location

CONTOUR INTERVAL 20 FEET

## EXPLANATION

<b>Quaternary</b>	<b>Qal</b> Alluvium - Stream deposits of unconsolidated to loosely consolidated clay- to boulder-sized clasts. Indicated by brackets on cross section A-A' where too thin to show. Estimated thickness up to 25 ft (7.6 m)	<b>Qc</b> Colluvium - Unconsolidated debris deposited by gravity along steep slopes. Indicated by brackets on cross section B-B' where too thin to show. Estimated thickness up to 20 ft (6.1 m)	<b>Qaf</b> Alluvial fan - Unconsolidated to loosely consolidated, clay- to boulder-sized clasts. Located proximal to present-day drainages	<b>Qt</b> Talus - Angular quartzite blocks derived from Xbq. Estimated thickness up to 33 ft (10 m)	<b>Qtd</b> Terrace deposit - Unconsolidated clay- to boulder-sized clasts deposited along present-day drainages. Thickness up to 35 ft (11 m)	<b>Qtr</b> Terrace deposit - Clay- to boulder-sized clasts of older stream deposits. Approximately 40-50 ft (12-15 m) above present-day drainages	<b>Qdf</b> Debris flow deposit - Composed of very large boulders within finer debris. Formed from breached dam(s) of landslides in lower Little Elk Canyon. Boulders have been washed onto terrace deposits near Interstate 90. Indicated by brackets on cross section A-A' where too thin to show. Thickness up to 40 ft (12 m)	
<b>Tertiary</b>	<b>Tg</b> Gravel deposit - Generally gravel at surface; artificial exposures may disclose pinkish-bentonitic beds at depth. Based on elevation, interpreted to be equivalent to known Oligocene White River deposits in other areas. Thickness up to 30 ft (9.1 m)	<i>Discontinuity</i>						
<b>Cretaceous</b>	<b>Ki</b> Lakota Formation - Sandstone, tan, thick-bedded. Only basal portion present. Exposed thickness 30-45 ft (9-14 m)	<i>Discontinuity</i>						
<b>Jurassic</b>	<b>Jmu</b> Morrison Formation and Unkapa Sandstone - Greenish-gray bentonitic shale and sandstone, and beige to white, friable dune sandstone. Combined thickness 100-125 ft (30-38 m)	<b>Jsu</b> Sundance Formation - Greenish-gray shale, brownish-orange sandstone, and minor limestone, laminated to medium-bedded. Thickness 330-350 ft (101-107 m)	<i>Discontinuity</i>				<b>Jm</b> Minnekahta Limestone - Pink to beige, thin-bedded limestone. Thin shale partings near the middle. Contains abundant stylonites. Thickness 38-42 ft (11-13 m)	
<b>Triassic</b>	<b>Tsp</b> Spearfish Formation - Red siltstone, shale, and minor limestone and gypsum. May include the Gypsum Spring Formation at the top of the unit. Thickness 500-530 ft (152-162 m)	<b>Tpm</b> Minnekahta Limestone - Pink to beige, thin-bedded limestone. Thin shale partings near the middle. Contains abundant stylonites. Thickness 38-42 ft (11-13 m)	<i>Discontinuity</i>				<b>Tpo</b> Opache Shale - Maroon to red shale and siltstone. Thickness 80-110 ft (24-34 m)	
<b>Permian</b>	<b>Pp</b> Minnelusa Formation - Brown, reddish, thin- to thick-bedded sandstone, limestone, and shale. Thickness 565-645 ft (172-197 m)	<i>Discontinuity</i>						
<b>Pennsylvanian</b>	<b>Pp</b> Minnelusa Formation - Brown, reddish, thin- to thick-bedded sandstone, limestone, and shale. Thickness 565-645 ft (172-197 m)	<i>Discontinuity</i>						
<b>Mississippian</b>	<b>Mp</b> Pahasapa Limestone - Beige to white, medium- to thick-bedded limestone and dolomite. Thickness 480-510 ft (146-155 m)	<i>Discontinuity</i>						
<b>Devonian</b>	<b>Mde</b> Englewood Limestone - Massive to pink, thin-bedded, finely crystalline shale, limestone, and dolomite. Thickness 36-55 ft (11-17 m)	<i>Discontinuity</i>						
<b>Ordovician</b>	<b>Ow</b> Winnipeg Formation - Gray to green, laminated to thin-bedded shale. Thickness 10-60 ft (3-24 m). Exposures north of Little Elk Canyon may include up to 5 ft (1.5 m) of Whitehouse Limestone	<i>Discontinuity</i>						
<b>Cambrian</b>	<b>Ocd</b> Deadwood Formation - Glaucous sandstone, shale, limestone pebble conglomerate, and local basal conglomerate. Uppermost thick-bedded iron-stained sandstone from 55-90 ft (17-27 m) thick. Total thickness 300-425 ft (107-130 m)	<i>Discontinuity</i>						
<b>Lower Proterozoic</b>	<b>Xmg</b> Metagabbro - Sills and dikes, greenish-gray, medium-grained except where sheared or well foliated. It is uncertain that all bodies shown as Xmg are of the same age	<b>Xmgo</b> Metagabbro (older) - Sills and dikes of greenish, massive to schistose amphibolite. Sill 0.75 mi (1.2 km) southeast of Green Mountain has a Pb-Pb zircon age of 1.964±0.15 Ma (Redden and others, 1990) indicating a post Estes Formation emplacement. It is uncertain that all bodies shown as Xmgo are of the same age	<b>Xbc</b> Hay Creek Greenstone - Xbc - Tholeiitic, generally pillowed greenstone and amphibolite. Xbcf - Interflow intervals of phyllite and metachert. May be locally similar to carbonate-facies banded iron-formation. Only thicker intervals shown	<b>Xf</b> Ferruginous metachert - Massive to thin-bedded, iron-stained metachert, typically interbedded with lenses of dark phyllite. Locally transitional to carbonate facies iron-formation	<b>Xbz</b> Buck Mountain Quartzite - Xbm - Phyllite and quartzite; thin- to thick-bedded, gray to olive-drab phyllite and tan to grayish to white quartzite; abundant ripple marks in metachert protomylonite. Xbz - Thick-bedded to massive quartzite subunits. Xbp - Dominantly thin-bedded greenish to gray phyllite with sparse quartzite. Protolith of the Buck Mountain Quartzite is a marine shelf deposit	<b>Xgd</b> Gingress Draw Slate - Olive-drab to tan, thin-bedded slate. Contains quartzite and metachert beds to the south in the Pactola Dam quadrangle and is very difficult to distinguish from parts of the Buck Mountain Quartzite. Chlorite-rich phyllite is likely reworked volcanic tuff	<b>Xrd</b> Roberts Draw Limestone - Xrp - Greenish-gray to dark-gray phyllite; locally contains 5 to 25 percent dolomite. Some magnetite occurs where adjacent to carbonate rocks. Very sparse small stromatolites locally occur in the more pure dolomite units. Xrd - Recrystallized siliceous dolomite; typically thick-bedded and locally containing lenses of pebbly metaconglomerate and quartzite	<b>Xes</b> Estes Formation - Xec - Metaconglomerate and impure quartzite. Cross pattern indicates areas containing clasts of banded iron-formation which postdates the Boxelder Creek Formation. Xeq - Quartzite; thin- to thick-bedded, gray, tan, grayish-white; may include gray phyllite and lenses of metaconglomerate. Xep - Gray phyllite. Xea - Arkosic quartzite and metagrit with local phyllite interbeds. Protolith is generally a marine fan deposit adjacent to growth faults. Approximately 2.10-2.02 Ga (Frei and others, 2008)
<b>Precambrian</b>	<b>Xbd</b> Blue Draw Metagabbro - Gravity differentiated sill consisting of a mafic lower cumulate zone (altered to serpentine) and a metagabbro interior. An upper, thin granitic differentiate is exposed in the Nemo quadrangle where a Pb-Pb age of 2.400±0.1 Ma was obtained (Dahl and others, 2006). Separate small body of serpentine along upper Jim Creek inferred to be correlative	<b>Xboc</b> Boxelder Creek Quartzite; includes Novak and Tomahawk tongues - Xbcq - Quartzite, grayish-white to tan. Thick-bedded with abundant small scale cross bedding indicative of a fluvial origin. Novak Tongue includes Xbnp, Xbnd, Xbnc, Xbnc, and Xbnc. Xbnp - Phyllite, greenish-gray, contains sparse quartzite and metachert clasts probably of Archean age which decrease in abundance to the north. Xbnd - Dolomite and minor gray phyllite. Xbnc - Metaprecambrian and quartzite. Abundant taconite clasts of hematite and metachert iron-formation typically decrease to the north and paraconglomerate intertongues with Xbnp. Xbnc - Metaconglomerate containing taconite clasts. Xbnc - Chlorite quartzite; locally has scattered pebbles and thin phyllite interbeds. Tomahawk Tongue includes Xbnc, Xbc, and Xbp. Xbc - Quartzite and metagrit. Xbc - Metaconglomerate and metagrit; typically a pyritic, graniferous, and surface pebble conglomerate. Granules of blue quartz probably derived from the Little Elk Granite are common. Usually has chromite and fuchsite. Xbp - Phyllite, gray to green, thin-bedded, poorly exposed. Protolith is an overbank deposit.	<b>XWu</b> Undifferentiated Lower Proterozoic and Upper Archean rocks - Shown only in cross sections	<i>Discontinuity</i>	<b>Wni</b> Nemo Iron-Formation - Wni - Thin-bedded metachert and hematite banded iron-formation. Locally consists of two or more subunits 15-56 ft (5-17 m) thick separated by similar thicknesses of thin-bedded greenish phyllite typically containing magnetite. Wnp - Green, thin- to medium-bedded phyllite; largely inferred but present in drill holes adjacent to Wni. No exposures known north of unit Wni. In fault contact with Proterozoic rocks but believed to be the source of banded iron-formation clasts in lower Boxelder Creek Quartzite. Approximately 2.9-2.6 Ga (Frei and others, 2008)	<i>Discontinuity?</i>	<b>W</b> Little Elk Granite - Generally coarse-grained, foliated, feldspar granitic which locally contains blue quartz. U-Pb age 2.549±0.11 Ma (Gosselin and others, 1988)	<b>Wu</b> Undifferentiated Upper Archean rocks - Shown only in cross sections

Contact  
Long dashed where approximately located; short dashed and queried where inferred on cross section; dotted where concealed

**FAULTS**  
Fault  
Long dashed where approximately located; short dashed where inferred; dotted where concealed; queried where uncertain  
Arrow indicates lateral movement. Tc indicates dip direction and amount. A (away) and T (toward) indicate relative movement on cross section. The E (east) fault is interpreted as a growth fault with erosion contemporaneous with lifting, infilling of the E (east) sediment to the west of the fault, and is characteristic of a rift environment

**FOLDS**  
Anticline  
Long dashed where approximately located; short dashed where inferred; dotted where concealed; queried where uncertain  
Syncline  
Showing trough line and direction of plunge. Long dashed where approximately located; short dashed where inferred; dotted where concealed; queried where uncertain  
Monocline, anticlinal bend  
Shorter arrow indicates steeper beds. Long dashed where approximately located; dotted where concealed  
Monocline, synclinal bend  
Shorter arrow indicates steeper beds. Long dashed where approximately located; dotted where concealed  
Minor fold  
Showing bearing and plunge

**STRIKE AND DIP OF BEDDING**  
Horizontal  
Inclined  
Top direction of beds known to be in dip direction  
Vertical  
Vertical fall indicates top direction of beds  
Overturned  
Top direction of beds known to be opposite dip direction  
Top of bed  
Shown by sedimentary structures

**STRIKE AND DIP OF FOLIATION**  
Inclined  
Vertical

**STRIKE AND DIP OF FRACTURES**  
Inclined  
Vertical

**KARST FEATURES**  
Stream loss zone

**OTHER FEATURES**  
Magnetic high  
Located by ground magnetic survey. Queried where uncertain  
Open pit mine  
Mine adit or cave  
Prospect pit  
Group of prospect pits  
Mine shaft  
Open pit mine or quarry  
In 4:1 materials  
4:1 slope

Publication Date: May 22, 2009

### Sources of Information Utilized in Map Compilation

Dahl, P.S., Hamilton, M.A., Wooden, J.L., Foland, K.A., Frei, R., McCombs, J.A., and Holm, D.K., 2006. 2480 Ma mafic magmatism in the northern Black Hills, South Dakota: A new link connecting the Wyoming and Superior cratons. *Canadian Journal of Earth Science*, v. 43, p. 1579-1600.

DeWitt, E., Buscher, D.P., Wilson, A.B., and Johnson, T.M., 1988. Map of mines, prospects, and patented mining claims, and classification of mineral deposits in the Nemo 7 1/2 minute quadrangle and the western one-third of the Rapid City 7 1/2 minute quadrangle, Black Hills, South Dakota. U.S. Geological Survey Open-File Report 87-0281D.

Frei, R., Dahl, P.S., Duke, E.F., Frei, K.M., Hansen, T.R., Fransson, M.M., and Jensen, L.A., 2008. Trace element and isotopic characterization of Neoproterozoic iron formations in the Black Hills (South Dakota, USA): Assessment of chemical change during 2.9-1.9 Ga deposition bracketing the 2.4-2.2 Ga first rise of atmospheric oxygen. *Precambrian Research*, v. 162, p. 441-474.

Gosselin, D.C., Papke, J.J., Zartman, R.E., Peterman, Z.E., and Laul, J.C., 1988. Archean rocks of the Black Hills, South Dakota: Reworked basement from the southern extension of the Trans-Hudson orogen. *Geological Society of America Bulletin*, v. 100, p. 1244-1259.

Redden, J.A., Peterman, Z.E., Zartman, R.E., and DeWitt, E., 1990. U-Th-Pb geochronology and preliminary interpretation of Precambrian tectonic events in the Black Hills, South Dakota. In LeVary, J.F., and Stauffer, M.R., eds., *The Early Proterozoic Trans-Hudson orogen of North America*. Geological Association of Canada Special Paper 37, p. 229-251.

Strobel, M.L., Sawyer, J.F., and Rahn, P.H., 2000. Field trip road log - hydrogeology of the central Black Hills of South Dakota. In Strobel, M.L., Davis, A.D., Sawyer, J.F., Rahn, P.H., Webb, C.J., and Naus, C.A., eds., *Hydrology of the Black Hills*. Rapid City, S. Dak., South Dakota School of Mines and Technology Bulletin 20, p. 240-245.

